

## Homework set #3 (assigned 17 October, due 24 October)

**General comments.** There is a lot of cool stuff in these problems and when you are done you will have derived some fairly sophisticated understanding of planets and the Solar System based on relatively simple ideas. All of these problems represent the techniques and approaches that are currently the state-of-the-art in planetary science, perhaps simplified somewhat.

I suggest you read the entire problem set through before starting any of it. The equations I give you might be useful for more than one of these problems. I suggest you start early so that you can ask me questions if you get stuck somewhere. You might need to think about some problems two or three times before the best approach becomes evident. A great technique here will be to write down everything you know for a given problem, write down what you are trying to find out, draw a picture, and then take it from there. Also, you may find that you need to look up various values for this problem set (like the temperature of the Sun or the radius of Mars); you should feel free to do so. Lastly, note that the symbol  $\odot$  refers to the Sun, so  $M_{\odot}$  is the mass of the Sun. I also remind you that an AU is an astronomical unit and that 1 AU is equal to  $1.5 \times 10^{11}$  m.

1) Io is the most volcanically active body in the Solar System. It also has the highest heat flow from its surface. Its heat flow is approximately  $2.5 \text{ W/m}^2$ ; for comparison, the Earth's heat flow is around  $0.05 \text{ W/m}^2$  (globally averaged). A Watt (W) is the metric unit of energy flux (energy per second), and a Joule (J) is the metric unit of energy:  $1 \text{ W} = 1 \text{ Joule per second}$ . What is the total heat energy escaping from Io, per second? Your answer should be either in J/s or in W.

What is the volume of Io's "crust"? You can assume the "crust" has a thickness of 10 km. If the density of this material is around  $3000 \text{ kg/m}^3$ , what is the mass of Io's crust?

The *latent heat of melting* is the amount of energy needed to convert a solid substance at its melting temperature to a liquid at the same temperature. For typical volcanic rocks, the latent heat of melting is around  $5 \times 10^5 \text{ J/kg}$ , which means it takes  $5 \times 10^5 \text{ J}$  to turn 1 kg of volcanic rock from solid at its melting temperature to liquid at the same temperature. Assuming that providing the latent heat of melting is the only important time-limiting step in turning Io's rocks to liquid, how much time does it take to completely melt all of Io's crust? This is the resurfacing time for Io. How many craters would you expect to see on Io? How many have you seen on Io (we have seen at least a couple pictures of Io so far in this class, or you can easily find pictures of Io on the web). What do you think about all this – does it make sense and hold together?

Io gets most of its heat flow from tidal pulling and pushing as it orbits Jupiter. Europa also orbits Jupiter and gets quite a bit of tidal pulling. The heat flow on Europa is around 100 times less than that on Io. The latent heat of melting from solid ice to liquid water is  $3.4 \times 10^5 \text{ J/kg}$ . If, as some people have proposed, the uppermost 100 km of Europa is a thick water ice layer, what is the timescale to melt this layer? Do you expect Europa to have no, few, some, or many craters on its surface, and why? This surface/interior ocean, some people think, is the reservoir of water that life on Europa could survive in – but only if there is enough liquid water that persists for long enough. (Note: The density of ice is around  $1000 \text{ kg/m}^3$ .)

You might want to know how to calculate the volume of a spherical shell. The volume of a spherical shell is given by  $4\pi r^2 \Delta r$  where  $r$  is the radius of the sphere and  $\Delta r$  is the thickness of the shell. This equation is only valid when  $\Delta r$  is much, much smaller than  $r$ .

2) It has been suggested that the most or all of the Earth's water (around  $10^{18} \text{ m}^3$ ) is not primordial (i.e., from the time of formation) but rather was brought in later by comets. Take a typical comet: 50% water ice, 10 km radius, density around  $2000 \text{ kg/m}^3$ . How many of these typical comets are needed to bring in

the Earth's entire water supply of  $10^{21}$  kg? If this massive influx of comets occurred during Late Heavy Bombardment (from around 4.5–3.9 billion years ago), what was the cometary influx rate (comets/year)? How many years go by between comet impacts? Compare this to the current cometary influx rate, which is something like  $10^{-8}$  comets per year. How many years go by between comet impacts with the modern cometary influx rate?

Comets may contain, in addition to water ice, some amino acids. What is the total mass of amino acids that could have been deposited on Earth during Late Heavy Bombardment if we assume the above scenario is correct? You can assume that comets are 0.0003% (by mass) amino acids. We know that a pile of  $6 \times 10^{23}$  molecules of glycine (a typical amino acid) has a total mass of 75 g and that an amino acid might have a size around  $10^{-9}$  m. Spread uniformly over the surface of the Earth, how thick is this extraterrestrially-derived amino acid layer? Are we talking about a thick or thin layer of imported amino acids? Discuss briefly (a few sentences).

3) Atmospheric pressure is simply the weight (remember Newton's Laws) of the atmosphere pushing down divided by the area upon which the atmosphere is pushing. The metric units of pressure are Newtons/m<sup>2</sup> (also called a Pascal; you might also have heard of pounds per square inch [PSI], which is the English units equivalent).

The surface (atmospheric) pressure on Venus is around 90 bars (surface pressure on the Earth is 1 bar, by definition). Venus' atmosphere is 96% CO<sub>2</sub>, compared to 0.035% in the Earth's atmosphere. How much CO<sub>2</sub> is in the Earth's atmosphere? To simplify this calculation, you can assume that the Earth's atmosphere is uniformly dense and has a height of 10 km (this is the Earth's atmospheric *scale height*). The density of air is around 1 kg/m<sup>3</sup>. How much CO<sub>2</sub> is in Venus' atmosphere? The total amount of carbon in/on the Earth is roughly  $10^{20}$  kg; what percent is in the atmosphere? If Venus has the same total carbon inventory, what percent of Venus' total carbon is in the atmosphere? What can you say about Venus' rocks?

Something to keep in mind: not all of the mass of a CO<sub>2</sub> molecule is carbon.

4) The Initial Mass Function (IMF) describes how many stars of a given mass form. One theory about the IMF suggests that smaller mass stars are more likely to form than larger mass stars in this way: for 1 star with a given mass, there are 2.3 stars with one tenth that mass. If you assume that the nearest 100 stars have nice, friendly masses of either 10, 1, or 0.1 solar masses, how many of each mass are there in the nearest 100 stars? What percentage of the total mass does each group have?

If some wise, distant alien (let's say 1000 parsecs away) were observing us and these nearby stars, what would be the total luminosity of these 100 stars, and what percentage of the total luminosity would each of the three groups contribute?

You could make a table for this problem which lists all these things for the three groups of stars. You can leave your answers in terms of solar masses and solar luminosities. You should discuss your answer briefly.

For the purposes of this problem, you can assume that the following relationship for an individual star is true:

$$\left(\frac{L}{L_{\odot}}\right) = \left(\frac{M}{M_{\odot}}\right)^3 \quad (1)$$

5) We have talked about habitable zones in class. It is actually pretty straightforward to calculate the location of the habitable zone, both as a function of stellar temperature (that is, for different *stellar spectral types*)

and distance, and as a function of time in the Solar System (remember the faint early Sun paradox and the continuous habitable zone).

The temperature of a planet ( $T_{pl}$ ) is given by the following:

$$T_{pl}^4 = \frac{T_{\star}^4 R_{\star}^2 (1 - A)}{4a^2} \quad (2)$$

where  $T_{\star}$  is the temperature of the star heating the planet,  $R_{\star}$  is the radius of the star,  $A$  is the albedo (reflectivity) of the planet, and  $a$  is the orbital distance of the planet.

Calculate the location of the habitable zone as a function of stellar temperature and distance. You should be able to reproduce the figure we saw in lecture. However, of course I want you to actually do the calculations and not just copy the figure from the lecture. Make a plot (stellar temperature versus distance from the star) that shows the location of the habitable zone. Mark the stellar spectral types on your plot as well. You can assume a planetary albedo of 0.5 (that is, 50% reflective; the Earth's albedo is more like 28%, but that's okay). You can also assume that stellar radii and stellar temperatures are related in this way:

$$\left(\frac{R}{R_{\odot}}\right) = \left(\frac{T}{T_{\odot}}\right)^{1.3} \quad (3)$$

which is not really exactly true but is good enough for us to use in this problem ( $R_{\odot}$  and  $T_{\odot}$  are the radius and temperature of the Sun). You should also ignore all atmospheric effects. This is obviously a terrible assumption, as a planet in the habitable zone with no atmosphere isn't going to be much of a habitable planet, but it makes the problem much easier.

You can also calculate the evolution of the location of the habitable zone in our Solar System over time. The luminosity of the Sun as a function of time is given by this (Gough 1981):

$$L_{\odot}(t) = \left[1 + \frac{2}{5} \left(1 - \frac{t}{t_{\odot}}\right)\right]^{-1} \times (L_{\odot, today}) \quad (4)$$

where  $t$  is time (in billions of years) and  $t_{\odot}$  is the current age of the Sun, 4.5 billion years; and  $L_{\odot, today}$  is the luminosity of the Sun today ( $3.9 \times 10^{26}$  J/sec). Using this equation, make a plot (distance versus time) that shows the location of the habitable zone in our Solar System over time.

You might also find it useful to know that the luminosity of a star – the amount of energy the star emits per second – is given by this equation:

$$L_{\star} = 4\pi R_{\star}^2 \sigma T_{\star}^4 \quad (5)$$

where  $L_{\star}$  is the luminosity (units: Joules per second);  $R_{\star}$  is the radius of the star (units: meters); and  $T_{\star}$  is the temperature of the star (in Kelvins). Sigma ( $\sigma$ ) is the Stefan-Boltzmann constant ( $5.67 \times 10^{-8}$  J/m<sup>2</sup>/K<sup>4</sup>/second).

On your plot showing the evolution of our Solar System's habitable zone with time, mark (shade in) the location of the *continuous habitable zone*. Also mark the Earth. Briefly (one or two sentences) comment. How would including a planetary atmosphere affect your answer?